

Constraining Antarctic ice sheet stability during the Last Interglacial

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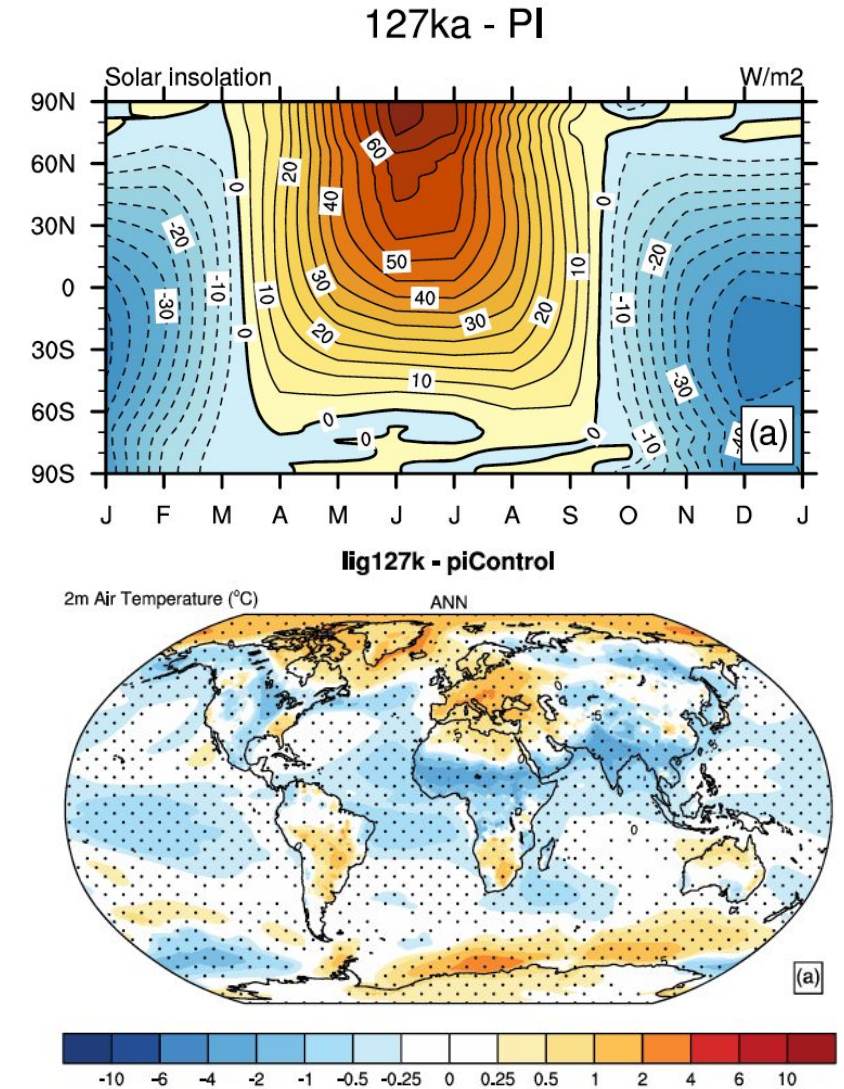


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BERN**

The Last Interglacial (LIG) (130,000 – 115,000 years ago)

1. Orbital configuration favored warm NH summers
 1. Perihelion was during Boreal Summer
 2. More eccentric orbit
 3. Higher obliquity
2. Much warmer Arctic than today
3. Lower CO₂ ~275 ppm



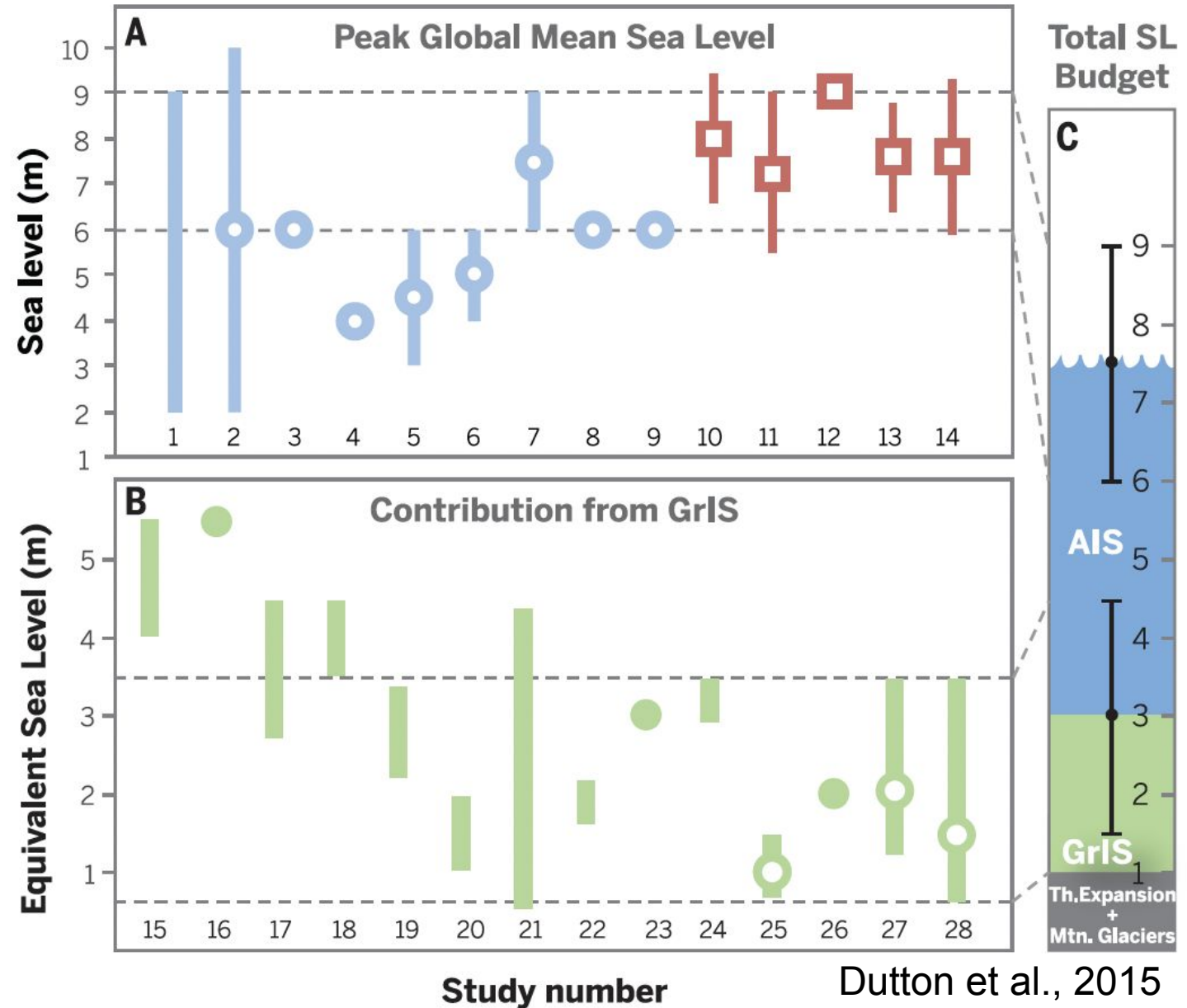
Otto-Bliesner et al., 2020

Last Interglacial sea level

Global mean sea level was
~**6-9 m** higher than present



Greenland contributed ~**1-3 m**
of sea level rise



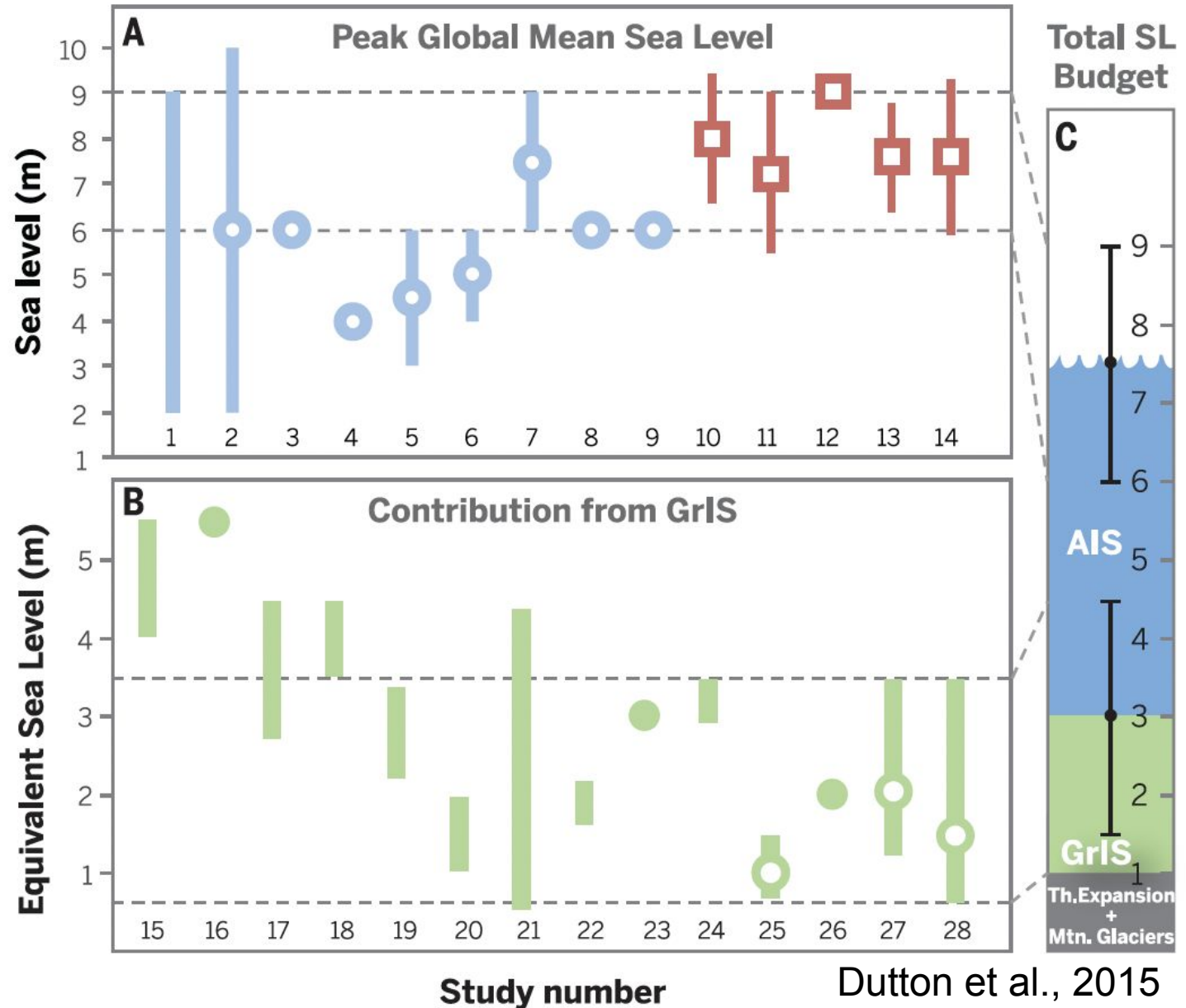
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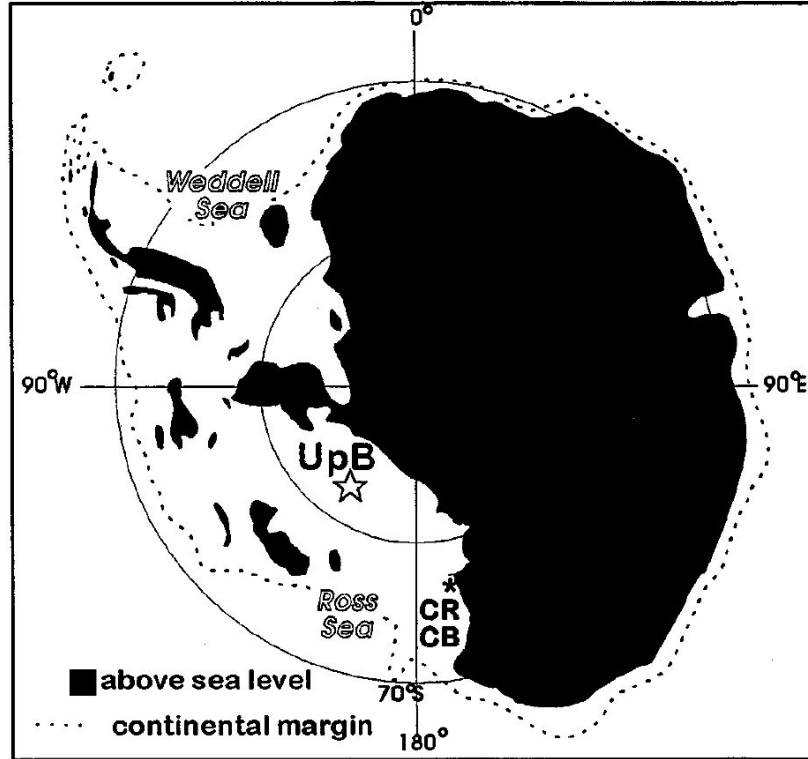


Greenland contributed ~**1-3 m**
of sea level rise

*Where does the
remaining ~3-6 m come
from?*

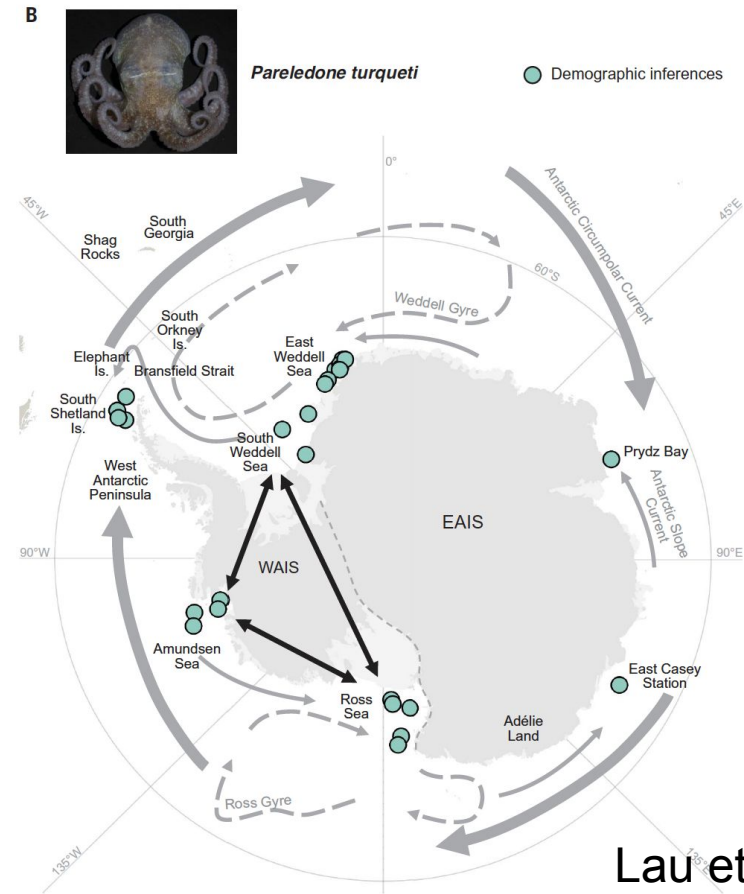


Evidence of a reduced Antarctic ice sheet during the LIG



Scherer et al, 1998

Late Pleistocene marine diatoms
under the West Antarctic Ice sheet.



Lau et al, 2023

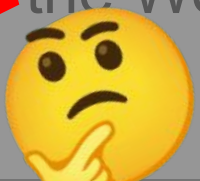
Genomic evidence of open ocean gateway
through the West Antarctic Ice Sheet during
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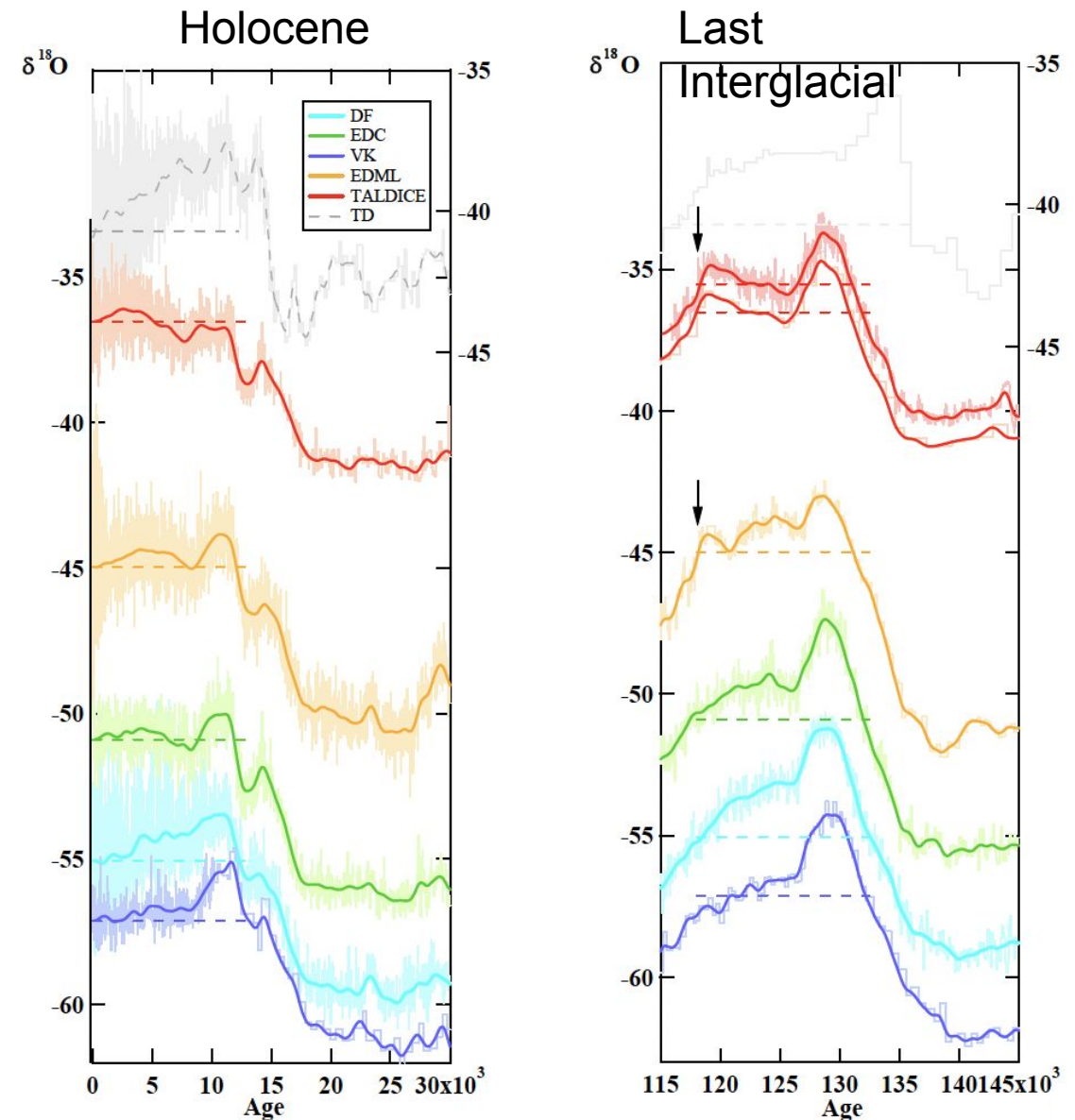
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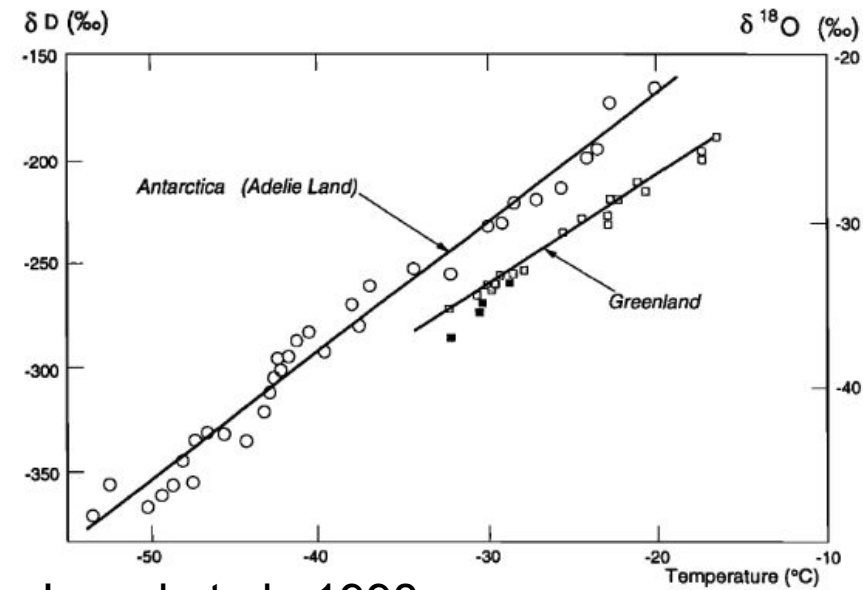
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Antarctic ice core records from the LIG

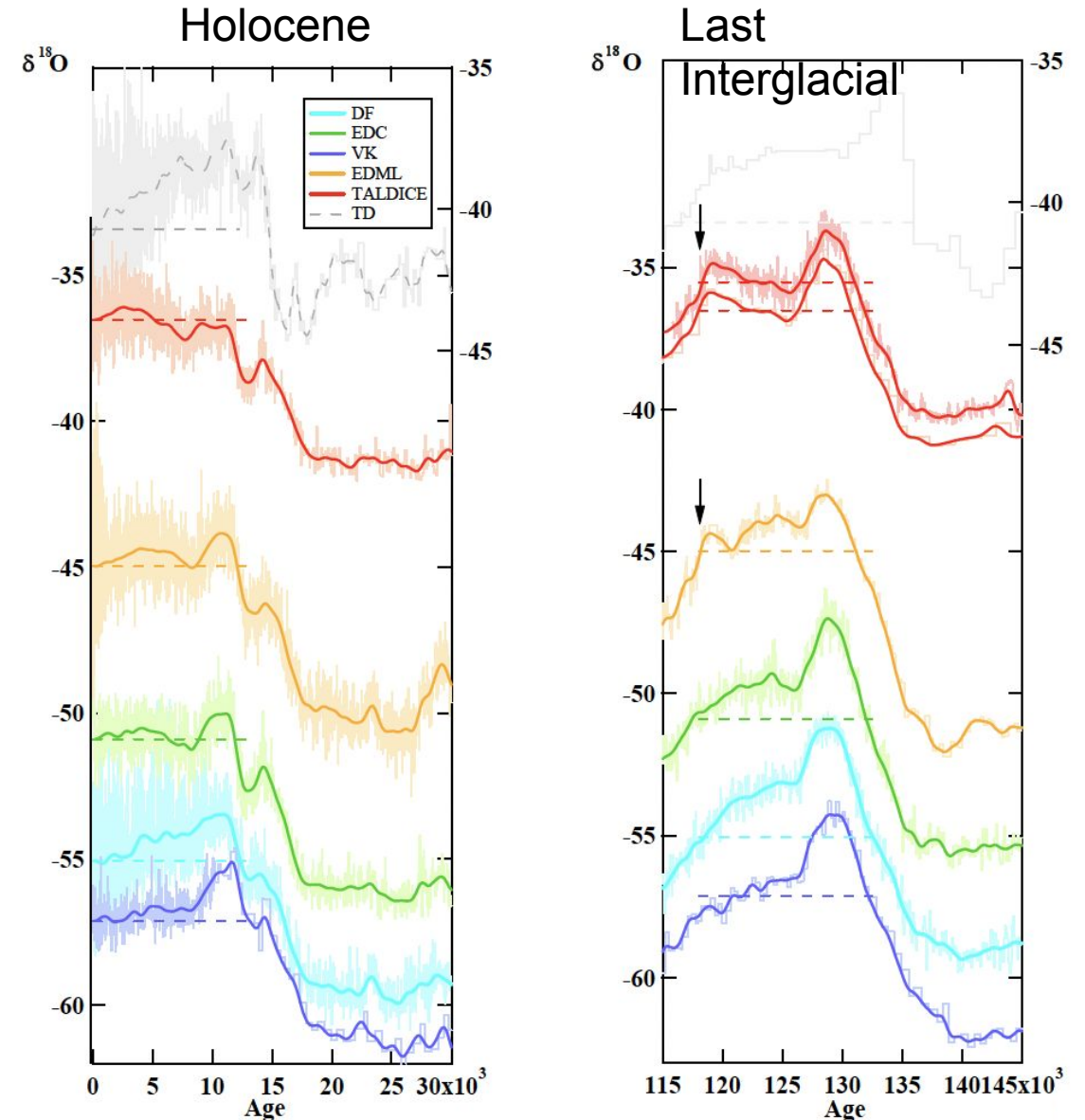


Antarctic ice core records from the LIG

Could Antarctic ice core $\delta^{18}\text{O}$ contain information about past ice sheet stability?



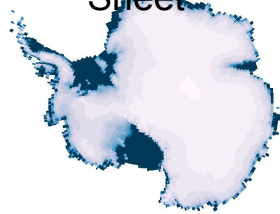
$\delta^{18}\text{O} \leftrightarrow \text{temperature} \leftrightarrow \text{background climate, elevation, etc.}$



LIG iCESM simulations

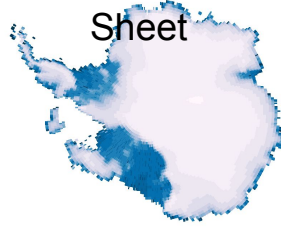
Ice Sheet
Configuration:

PI Antarctic Ice
Sheet



~0 m SLE

Minimally Collapsed
West Antarctic Ice
Sheet



~3 m SLE

Partially Collapsed
West Antarctic Ice
Sheet



~3 m SLE

Fully Collapsed
West Antarctic Ice
Sheet

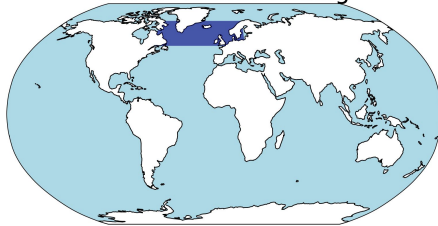


~6 m SLE

North Atlantic
Freshwater:

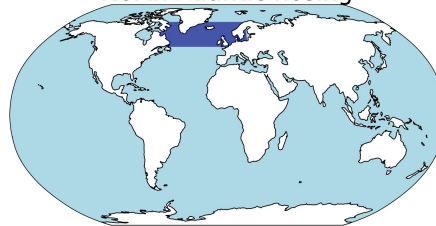
_NAFWF

North Atlantic hosing



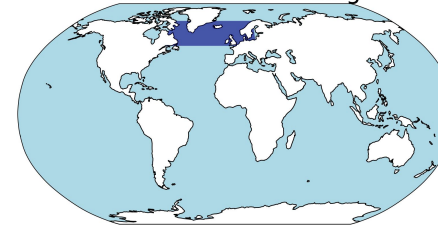
0.2 Sv x 1000
years

North Atlantic hosing



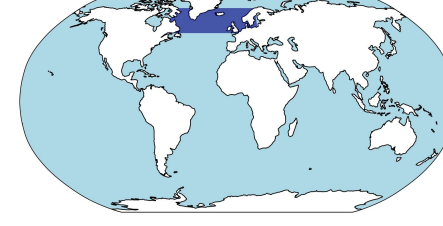
0.2 Sv x 1000
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North Atlantic hosing



0.2 Sv x 1000
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North Atlantic hosing

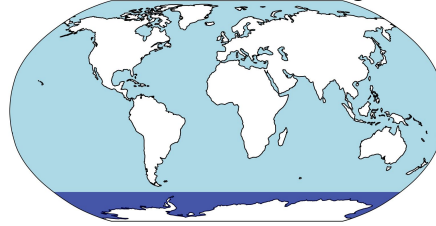


0.2 Sv x 1000
years

Southern Ocean
Freshwater:

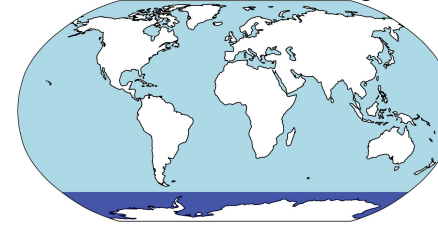
_SOFWF

Southern Ocean hosing



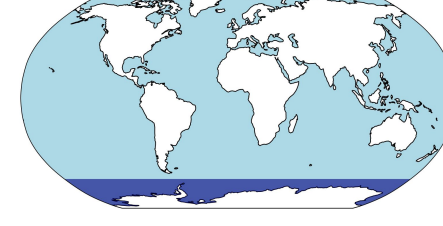
0.2 Sv x 500
years

Southern Ocean hosing



0.2 Sv x 500
years

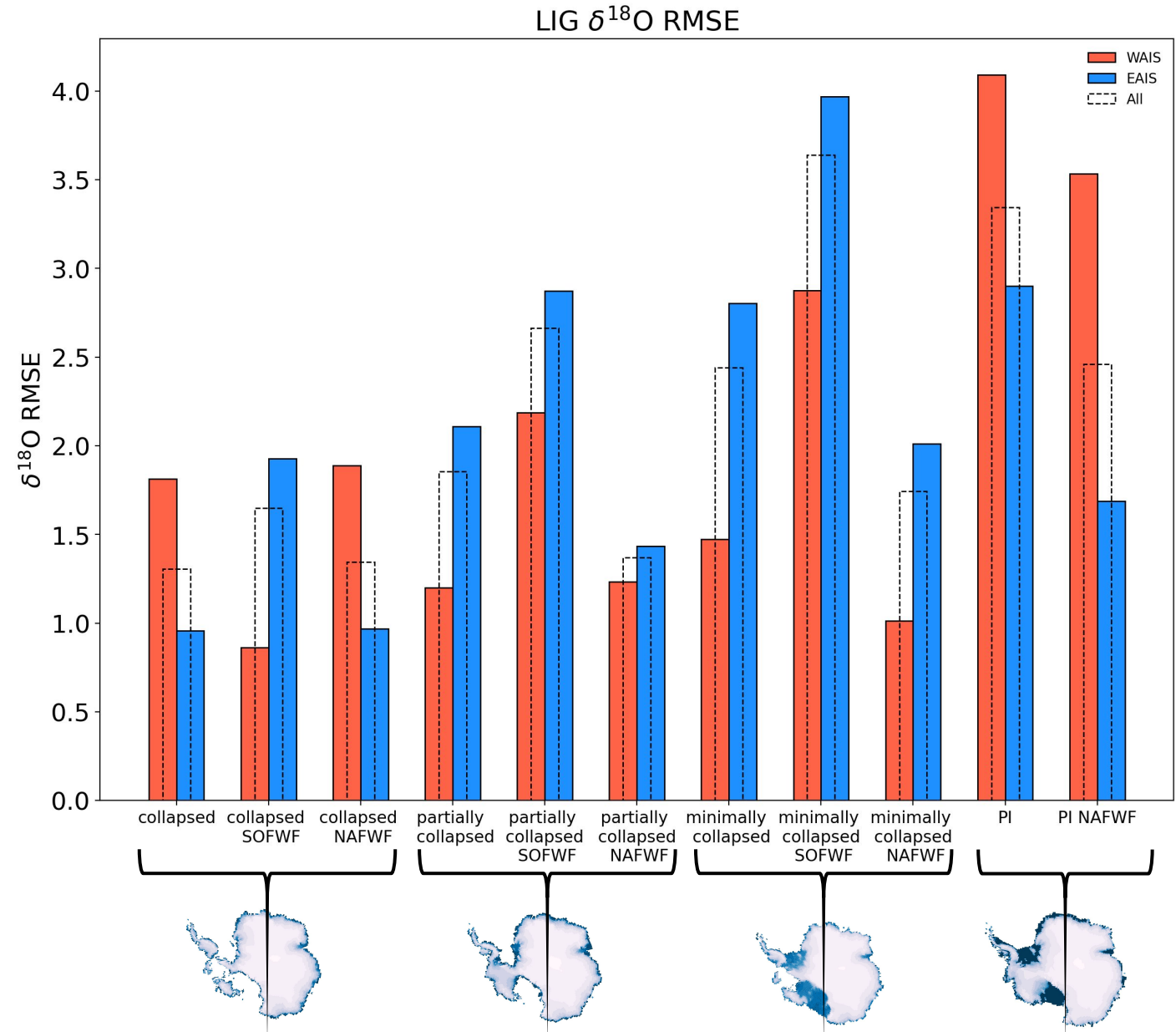
Southern Ocean hosing



0.2 Sv x 500
years

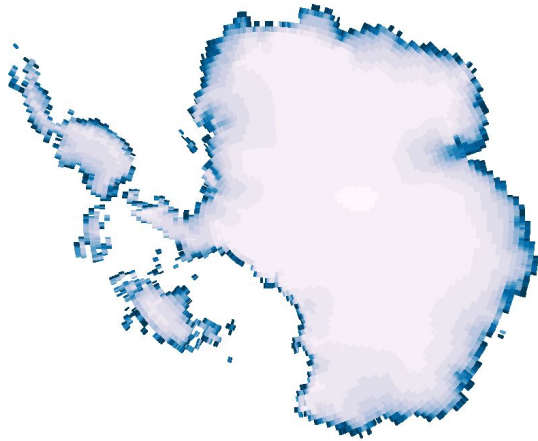
12 Simulations in total (11 LIG, 1 PI control)

Which simulation reproduces LIG $\delta^{18}\text{O}$ best?

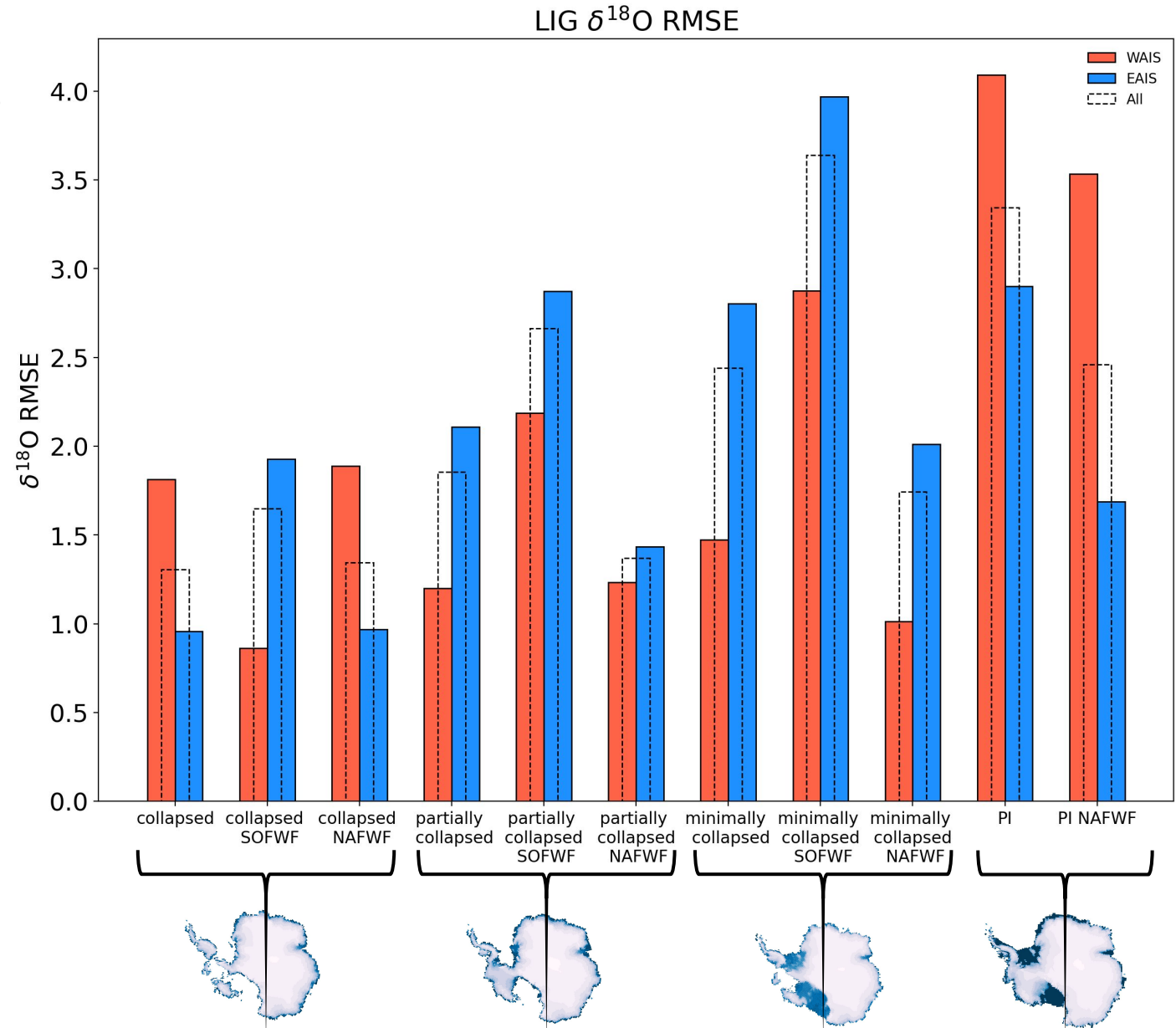


Which simulation reproduces LIG $\delta^{18}\text{O}$ best?

The fully collapsed Antarctic ice sheet configuration reduces RMSE the most.

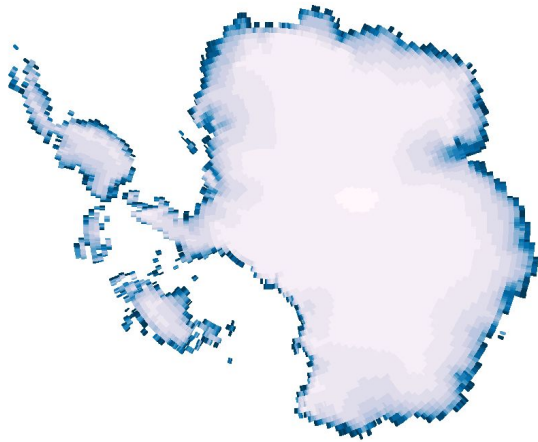


- NAFWF and SOFWF bracket $\delta^{18}\text{O}$ anomalies.
- Reality is somewhere in between?

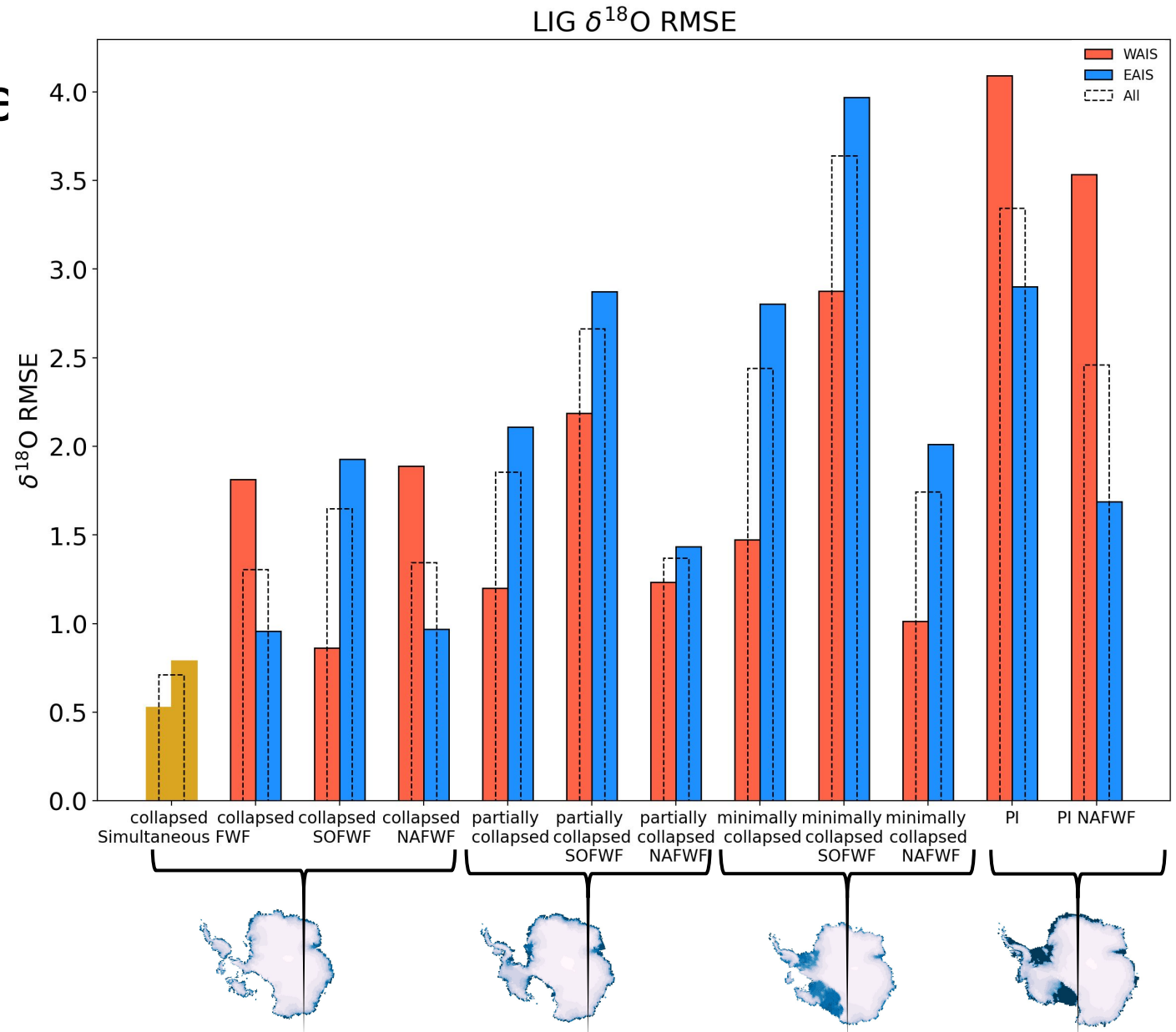


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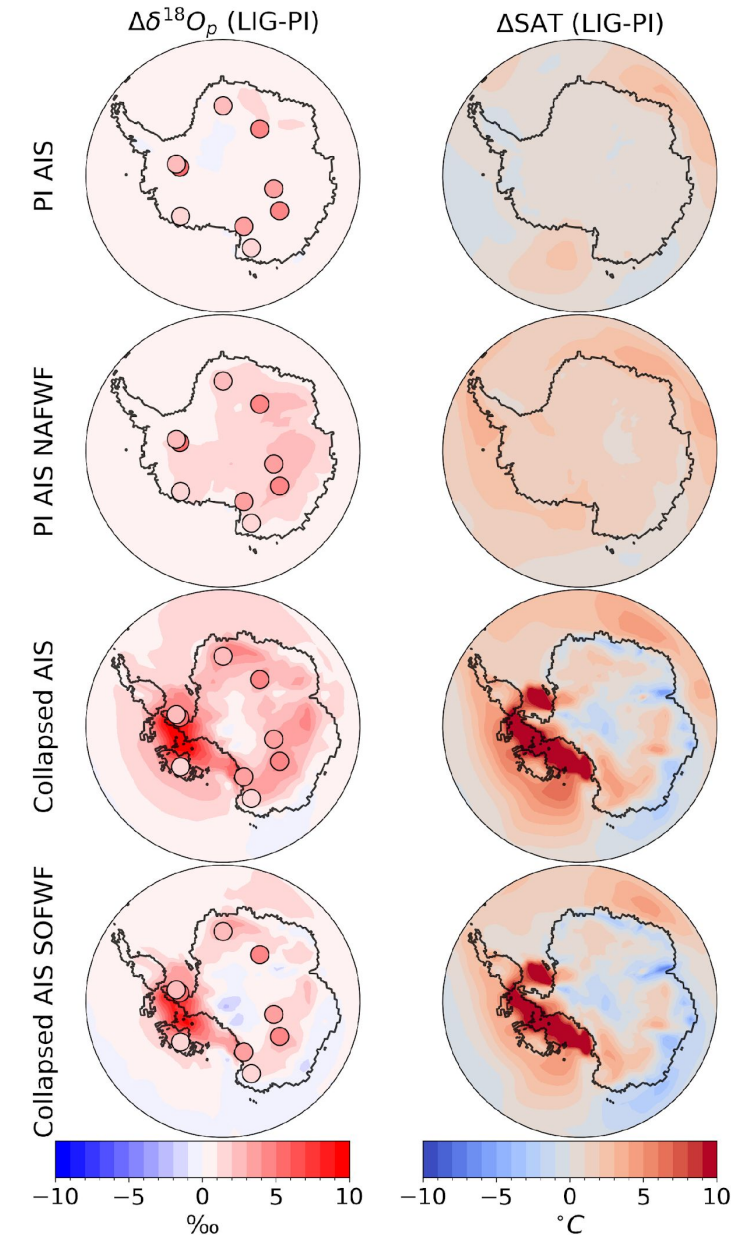


- NAFWF and SOFWF bracket $\delta^{18}\text{O}$ anomalies.
- Reality is somewhere in between?
- Applying the SOFWF impact to the collapsed WAIS with NAFWF gives lowest RMSE



Antarctic response to ice sheet collapse and hosing

- LIG orbit + GHGs cause some regional warming
- North Atlantic hosing and WAIS collapse lead to much more warming
- Southern Ocean hosing has a weaker impact on Antarctic surface temperatures than topographic change.

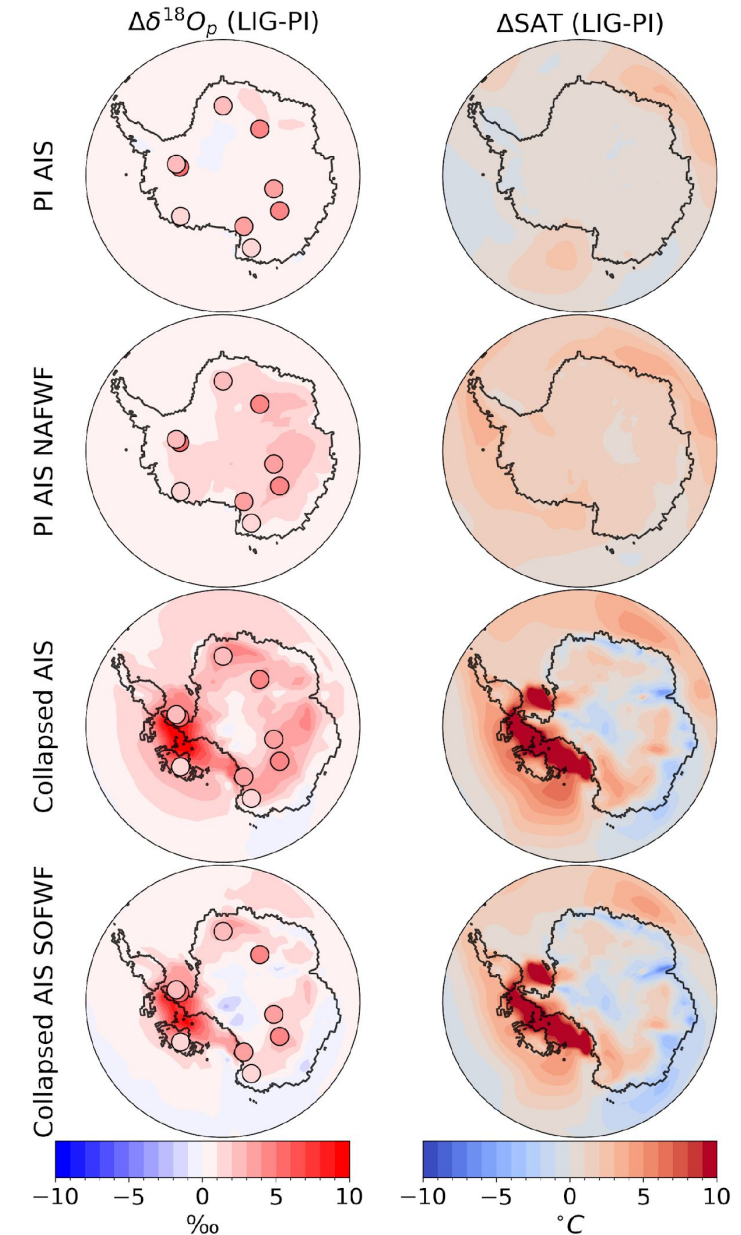


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Two potential controls on precipitation $\delta^{18}\text{O}$:

1. Lower Elevation ☐ Warmer condensation temperatures
☐ enriched $\delta^{18}\text{O}$
2. Reduced ice sheet ☐ regional warming ☐ more local precipitation
☐ enriched $\delta^{18}\text{O}$



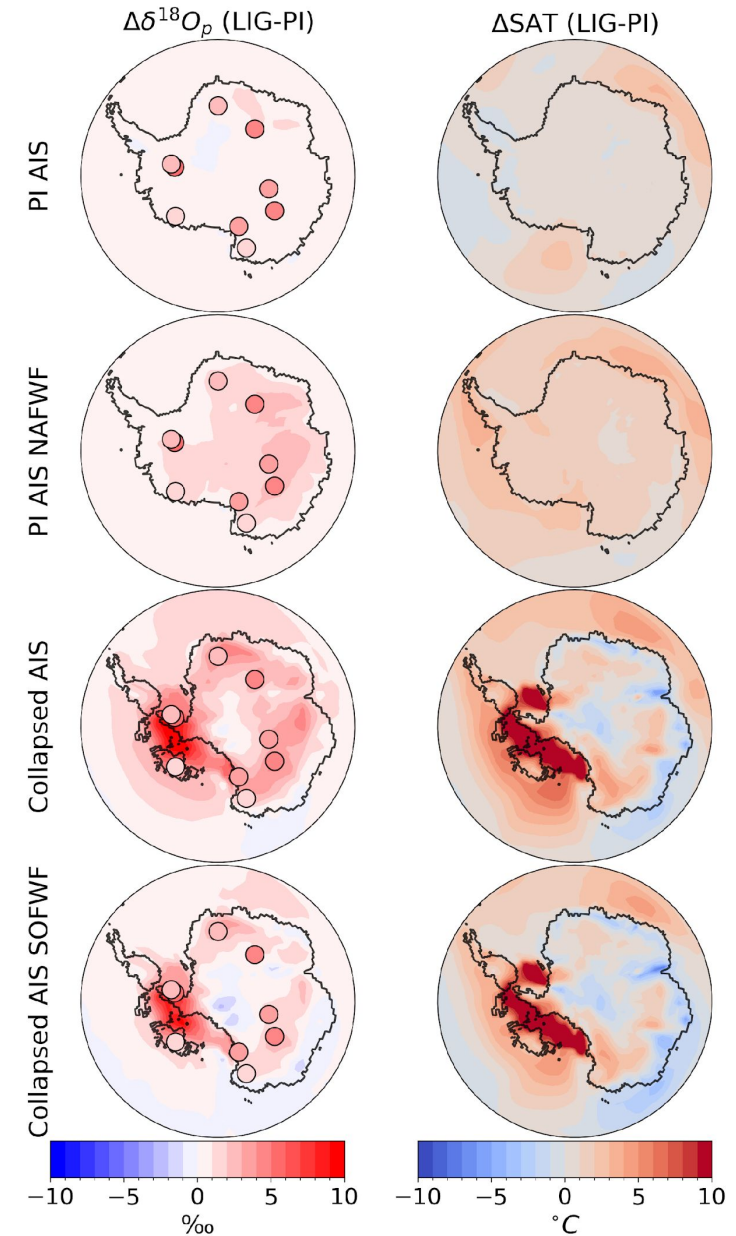
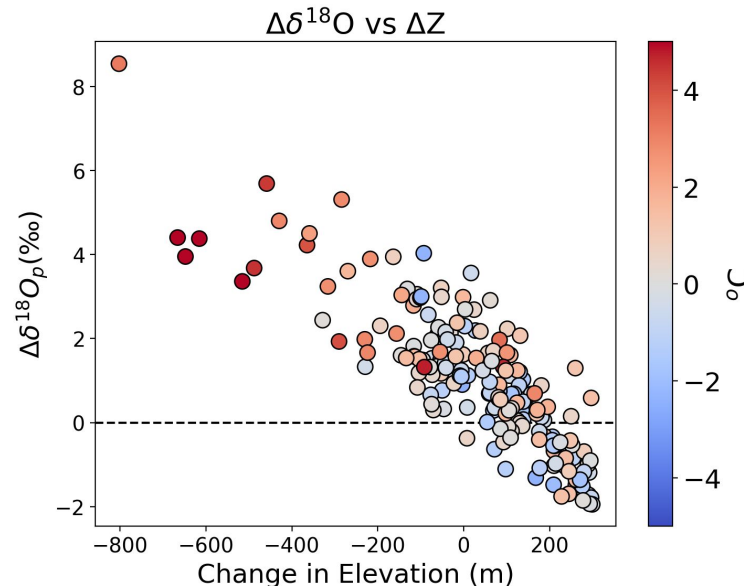
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$\delta^{18}\text{O}$ increases with decreasing elevation

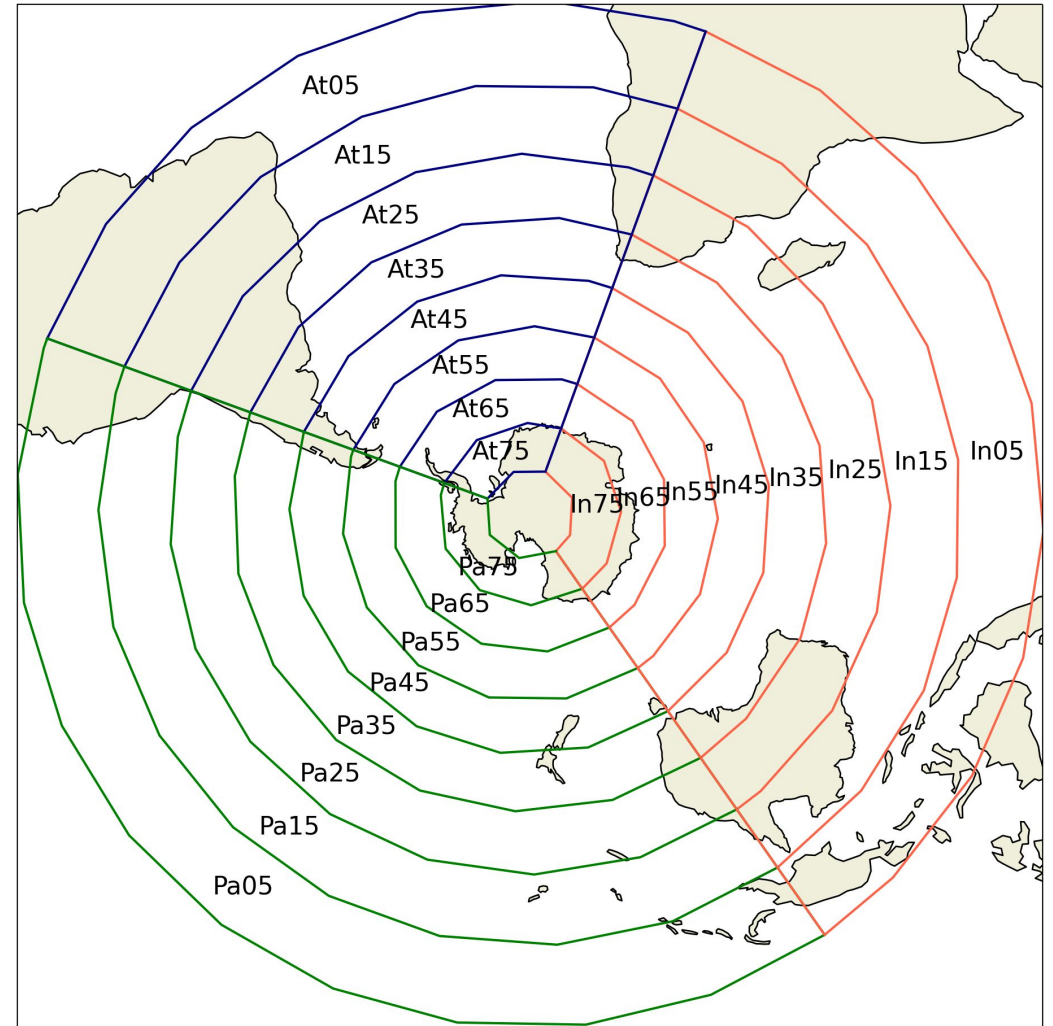


Adding water tags to our simulations

Tagging evaporative fluxes from specific regions allows us to track moisture sources and water isotope ratios from source to sink.

Water tags:

- Indian Ocean – 10° latitudinal steps (0-80°S)
- Pacific Ocean – 10° latitudinal steps (0-80°S)
- Atlantic Ocean – 10° latitudinal steps (0-80°S)
- Antarctic sourced moisture
- Southern Ocean sea ice sourced moisture



Moisture tagging is an extremely useful tool and should totally be added to CESM3. I know someone who would love to do that for a postdoc if y'all are hiring! 😊

Method for partitioning $\delta^{18}\text{O}$ signal

Following Hu et al., 2019:

$$\Delta(\delta^{18}\text{O}_P)_i \\ = \Delta[\textit{condensation}_i + \textit{rainout}_i + \textit{source composition}_i] \times \left(\frac{P_i}{P_{total}} \right) + \textit{source location}_i$$

- $\textit{condensation}_i = (\delta^{18}\text{O}_{P_{sink}} - \delta^{18}\text{O}_{wv_{sink}})_i$
- $\textit{rainout}_i = (\delta^{18}\text{O}_{wv_{sink}} - \delta^{18}\text{O}_{wv_{source}})_i$
- $\textit{source composition}_i = (\delta^{18}\text{O}_{wv_{source}})_i$
- $\textit{source location}_i = \delta^{18}\text{O}_{P_{sink}} \times \Delta \left(\frac{P_i}{P_{total}} \right)$

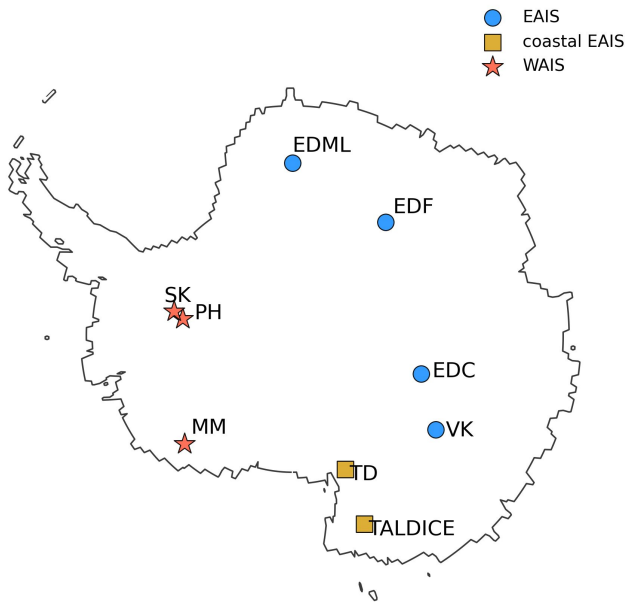
Core drivers of $\delta^{18}\text{O}$ change due to collapse

$$\text{condensation}_i = (\delta^{18}\text{O}_{P_{\text{sink}}} - \delta^{18}\text{O}_{wv_{\text{sink}}})_i$$

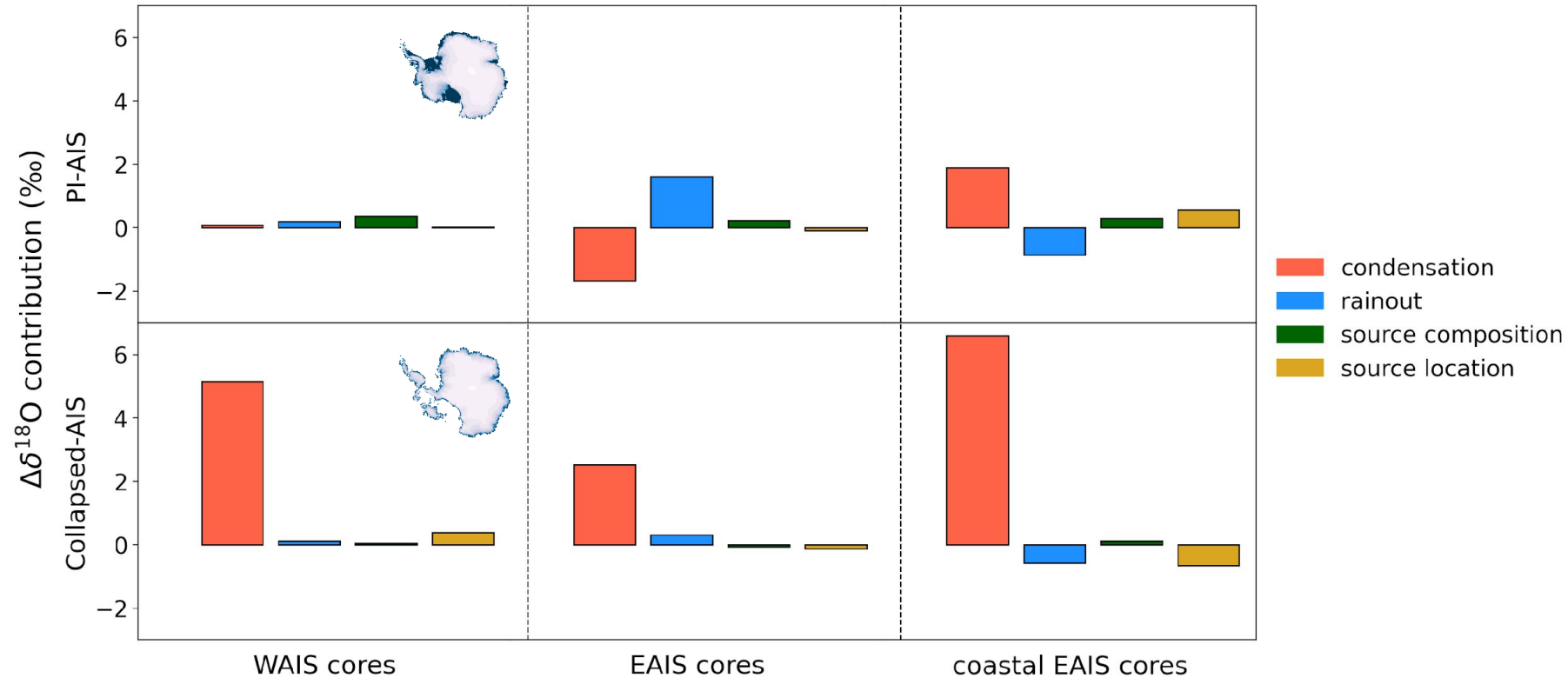
$$\text{rainout}_i = (\delta^{18}\text{O}_{wv_{\text{sink}}} - \delta^{18}\text{O}_{wv_{\text{source}}})_i$$

$$\text{source composition}_i = (\delta^{18}\text{O}_{wv_{\text{source}}})_i$$

$$\text{source location}_i = \delta^{18}\text{O}_{P_{\text{sink}}} \times \Delta \left(\frac{P_i}{P_{\text{total}}} \right)$$

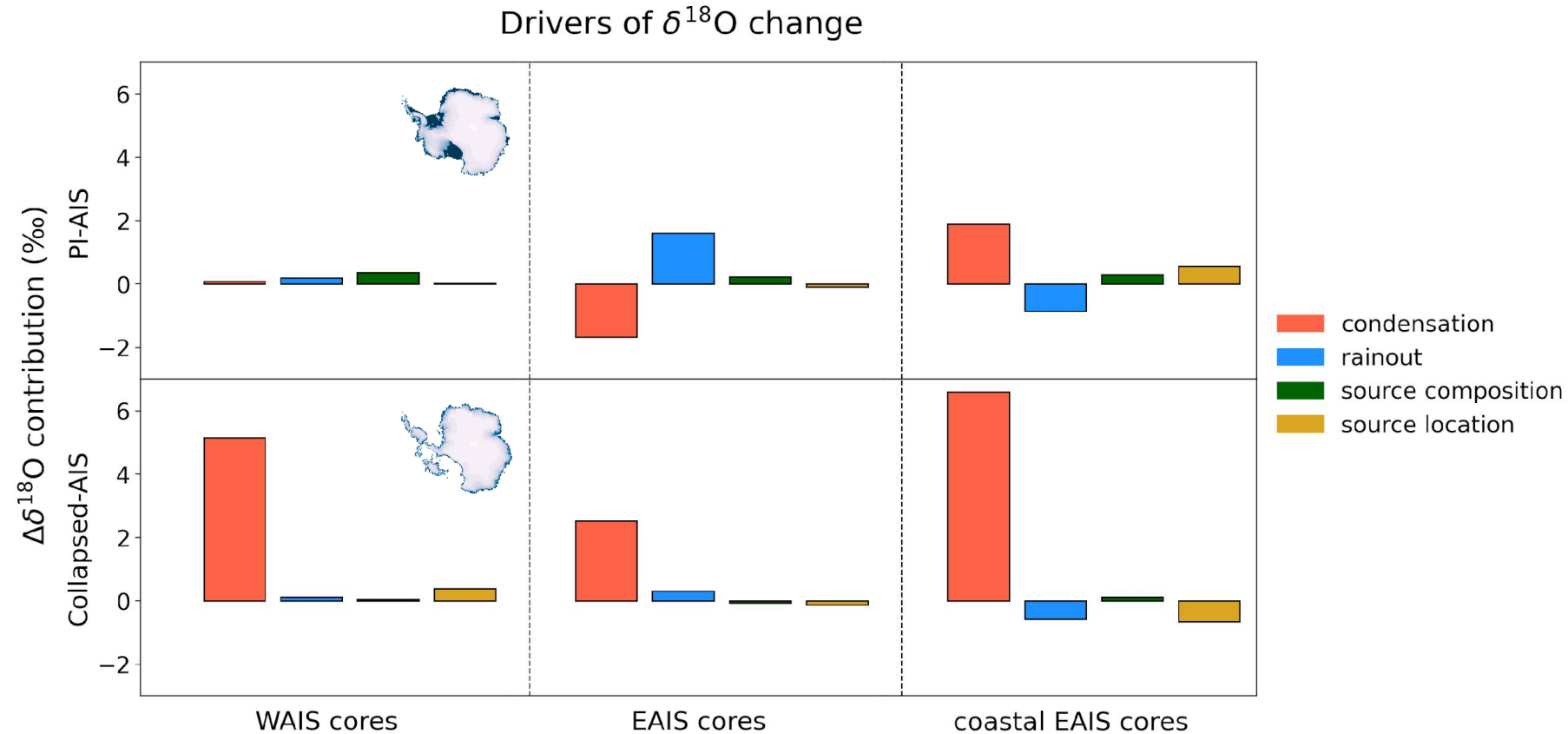
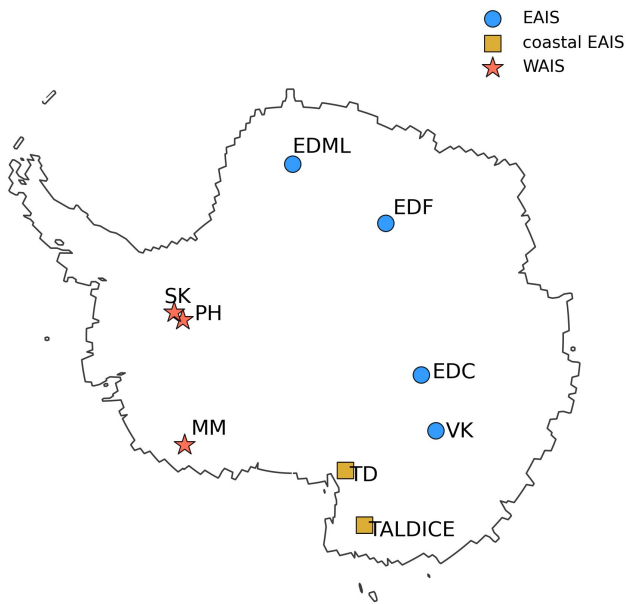


Drivers of $\delta^{18}\text{O}$ change



Core drivers of $\delta^{18}\text{O}$ change due to collapse

$$\begin{aligned} \text{condensation}_i &= (\delta^{18}\text{O}_{P_{\text{sink}}} - \delta^{18}\text{O}_{wv_{\text{sink}}})_i \\ \text{rainout}_i &= (\delta^{18}\text{O}_{wv_{\text{sink}}} - \delta^{18}\text{O}_{wv_{\text{source}}})_i \\ \text{source composition}_i &= (\delta^{18}\text{O}_{wv_{\text{source}}})_i \\ \text{source location}_i &= \delta^{18}\text{O}_{P_{\text{sink}}} \times \Delta \left(\frac{P_i}{P_{\text{total}}} \right) \end{aligned}$$



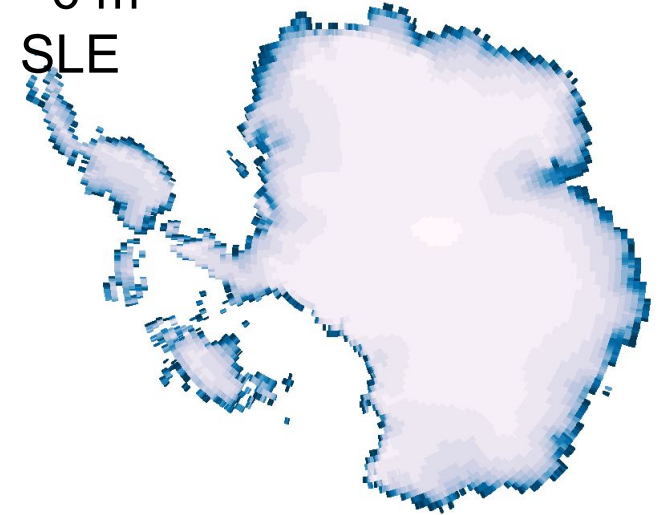
1. Condensation dominates $\delta^{18}\text{O}$ change in collapsed WAIS case
2. Aggregate impact of other drivers of $\delta^{18}\text{O}$ change is small

Thanks for listening! Questions? Feedback?

Recap:

1. A fully collapsed WAIS reproduces peak $\delta^{18}\text{O}$ best during the LIG
2. Collapsing the ice sheet drives regional warming, increasing condensation temperatures and enriching $\delta^{18}\text{O}$
3. Changing moisture sources and trajectories also contribute to $\delta^{18}\text{O}$ enrichment, especially over West Antarctica and peripheral East Antarctica. Could be evident in d-excess records.

~6 m
SLE



**Fully collapsed West
Antarctic Ice Sheet
127,000 years ago**

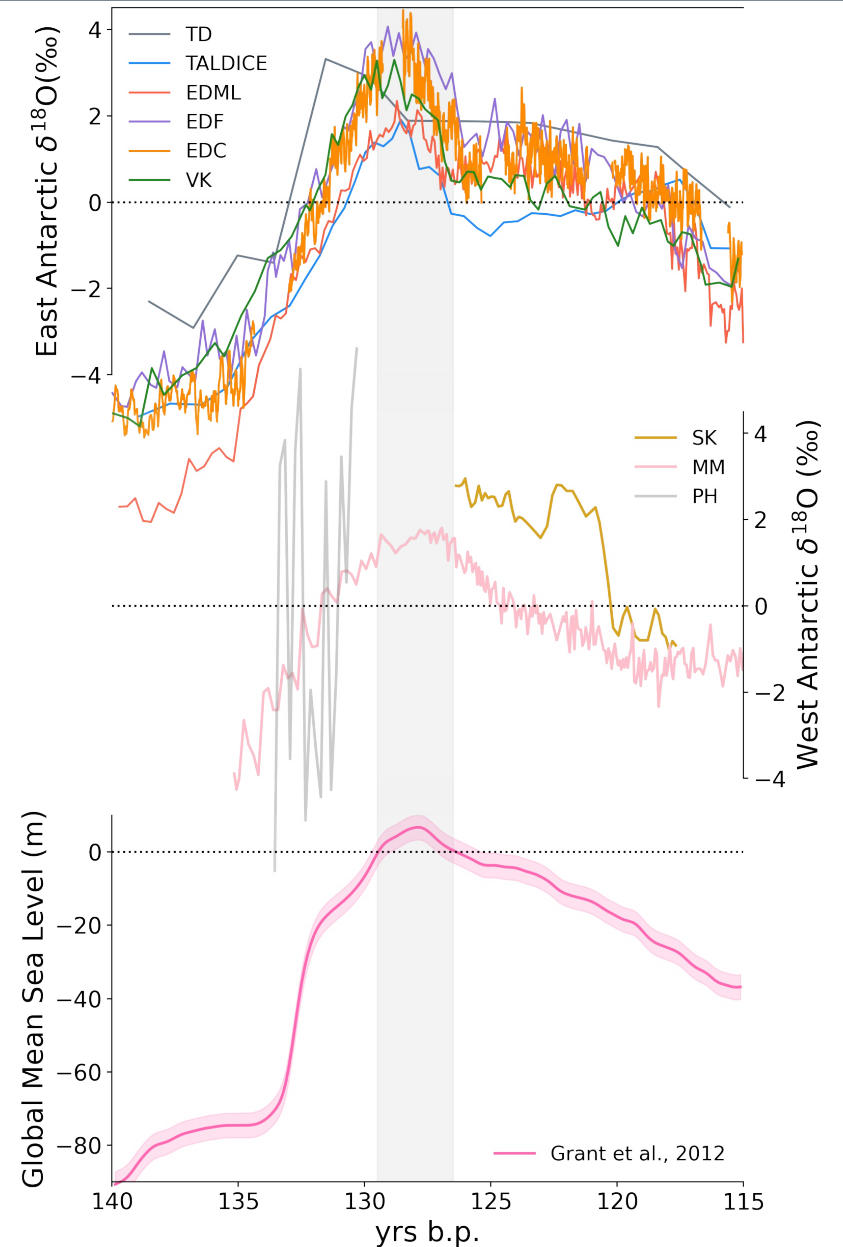
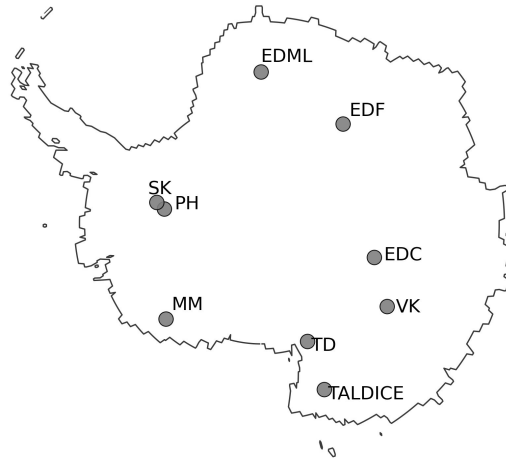
Email: joseph.schnaubelt@uconn.edu

Antarctic ice core records from the LIG

6 East Antarctic Ice Cores
(enriched 2-4 ‰)

3 West Antarctic Ice Cores
(enriched 2-5? ‰)

Synchronous with peak in
global mean sea level

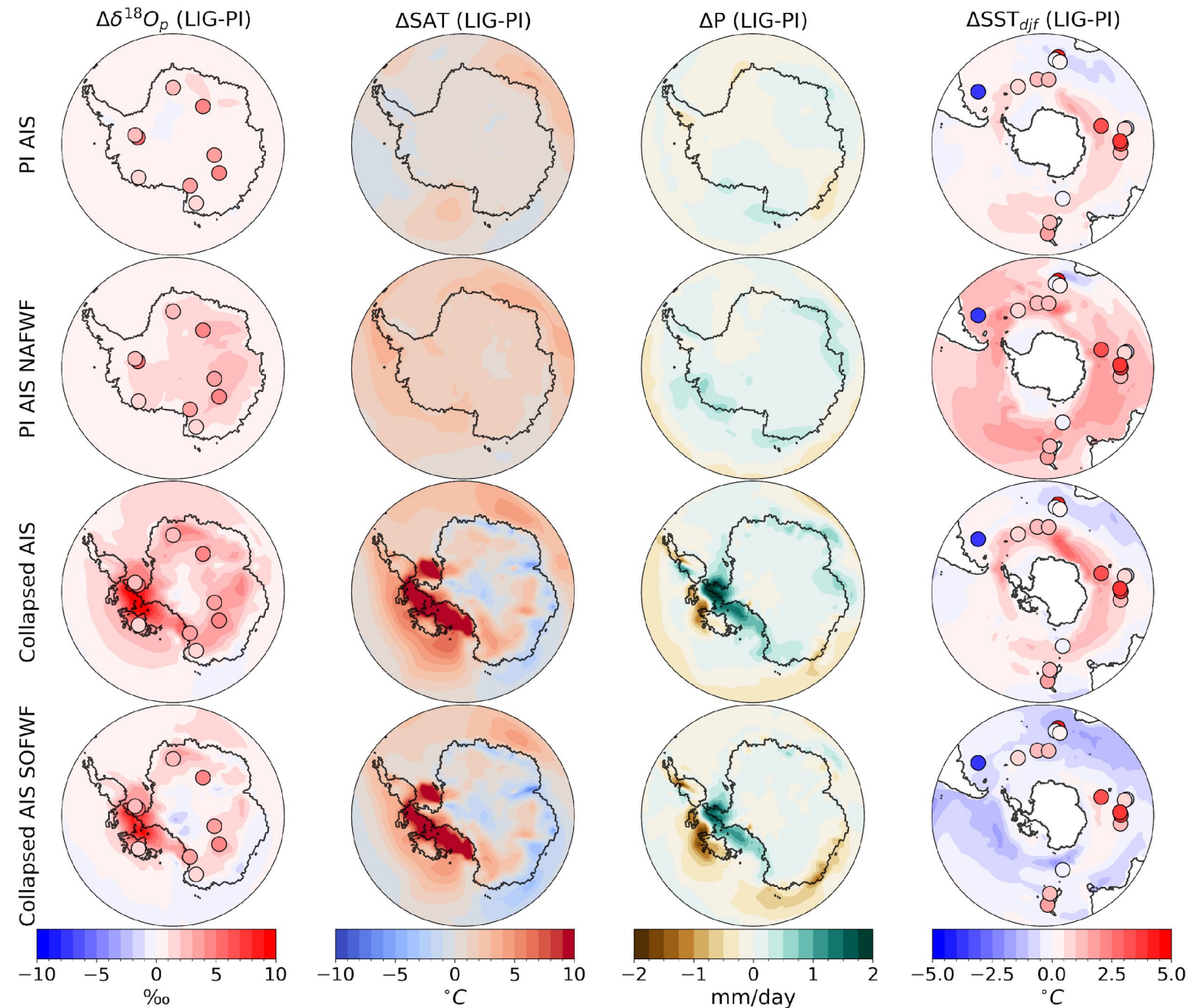


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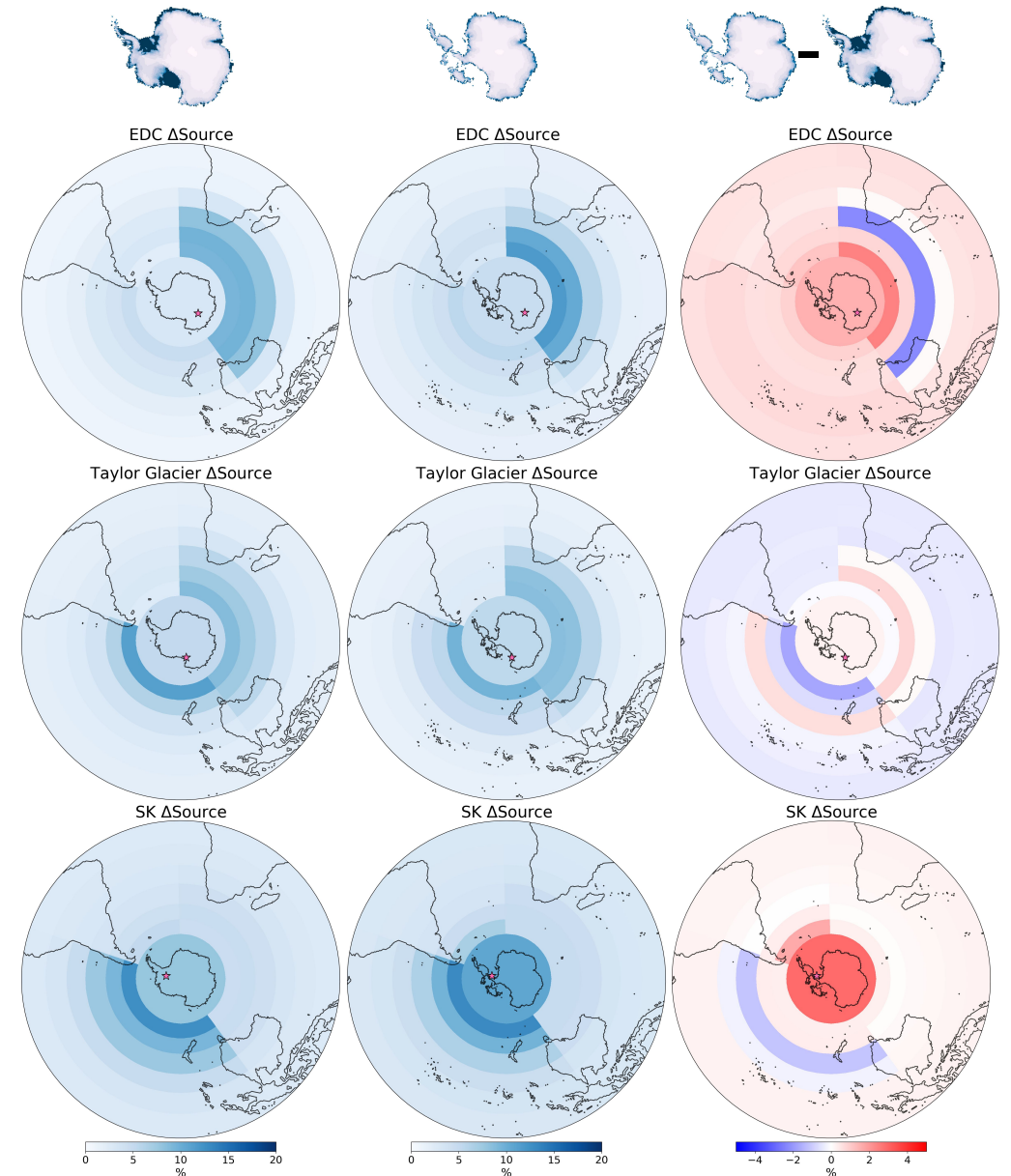


Antarctic moisture source change

Tracking the % contribution of moisture to 3 sample ice cores:

- Epica Dome C (EDC) - EAIS
- Skytrain Ice Rise (SK) - WAIS
- Taylor Glacier – coastal EAIS

1. EDC and SK experience a southward shift in moisture source



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1. EDC and SK experience a southward shift in moisture source
2. Taylor Glacier sees more northward moisture source.

Potentially due to an increase in coastal easterly wind strength.

Could be evident in d-excess records!

