Projecting climate responses using an AI Implementation of the Fluctuation-Dissipation Theorem trained on CESM2 Large Ensemble data

Haruki Hirasawa¹, Sookyung Kim², Salva Ruhling Cachay², Subhashis Hazarika², Peetak Mitra², Dipti Hingmire¹, Hansi Alice Singh¹, Kalai Ramea², and Phil Rasch³

¹University of Victoria, ²Palo Alto Research Center, ³University of Washington









Fluctuation Dissipation Theorem (FDT)

- The FDT states that the response of a dynamical system to a forcing can be estimated from the statistics of internal fluctuations.
- The climate is one such system where the FDT can be applied:



Fluctuation Dissipation Theorem (FDT)

 One common formulation of FDT for climate is in terms of a Linear Response Function L (e.g., Cionni et al., 2004, Gritsun and Branstator, 2007, Liu et al., 2018):

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Projected response Forcing

\delta \overline{y}(t) = L \delta f(t)
Linear Response Function

L = \int_{0}^{\infty} C(\tau)C(0)^{-1}d\tau
For covariance matrix C(\tau) at time lag \tau

and autocovariance matrix C(0)
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Limitations of classic FDT

- Applying FDT typically assumes:
 - 1. Dimension reduction
 - 2. Near-Gaussian statistics
 - 3. Linear responses
 - 4. Large reference internal variability datasets
- Can we loosen some of these requirements using AI?

Al Model Design

- We developed a spherical Multilayer Perceptron (MLP) model
- Maps monthly mean radiation anomalies to select climate variables at different time lags
- Data is regridded to spherical icosahedral grid to avoid dealing with grid area variations



CESM2 LE Training data

- Large amount of internal variability information is required to train this AI model
- We use data from the CESM2 smoothed biomass burning historical simulations
- Data is preprocessed by removing the ensemble mean for each grid point and month in the time series



Applying FDT with AI models

- We train the AI model on at a range of time lags au
- The AI models (A_{τ}) are regression models that map from an input field $\vec{x_i}$ to the average of possible trajectories of $\vec{y_i}$ for a given lag:

$$A_{\tau}(\overrightarrow{x_i}): \overrightarrow{x_i}(t) \rightarrow \overrightarrow{y_i}(t+\tau)$$

• To project a climate response $\langle \vec{y}(t) \rangle$, we integrate over the AI projections, which are the average of emulated responses to N different input fields $\vec{x_i}$:

$$\langle \vec{y}(t) \rangle = \sum_{\tau=0}^{T_{MAX}} \frac{1}{N} \sum_{i=0}^{N} (A_{\tau}(\vec{x}_i + \delta \vec{f}(t-\tau)) - A_{\tau}(\vec{x}_i))$$

• Where $\tau_{MAX} = 60$ months

Test case: Marine Cloud Brightening

- Marine Cloud Brightening (MCB) is a proposed solar radiation geoengineering technology
- Sea salt aerosol is injected into marine boundary layer clouds to increase their albedo
- Wide range of possible MCB forcing patterns, which are not practical to assess with ESMs



Regional Marine Cloud Brightening

- We impose MCB in CESM2 by setting cloud droplet number concentrations to 600cm⁻³ in three regions
- A constant $\delta \vec{f}$ for the AI model is calculated with radiative flux anomalies (ERF) from fixed-SST simulations

Perturbations applied to AI model



Spatial patterns of AI-FDT projections

- We compare the CESM2 "true" response (left) to the AI model response (right)
- The AI model overestimates the magnitude of the response, but correctly projects the pattern in most regions



Projected response to MCB in different regions

- Larger discrepancies between CESM2 and AI for individual regions
- Most of the total pattern is the result of La Nina-like response to SEP forcing
- Al expects stronger La Nina response to NEP than CESM2
- Non-ENSO signals are also captured (SEA cooling -> Amazon warming)



Conclusions

- We have developed an *ad hoc* implementation of FDT using an AI model to generate the response function
- The large data pool provided by the CESM2 LE is crucial for our AI model training
- Our AI-FDT model skilfully projects climate response **pattern** to MCB, but overestimates the magnitude of the response
- AI-FDT could be a useful tool for generating first look estimates of climate responses to forcing
 - E.g., for scenario development

Future Work

- 1. Uncertainty estimation:
 - Inter-ESM variability uncertainty : Train on different ESM Large Ensembles (SMILEs)
 - Inter-ESM forcing uncertainty : Apply ERF fields estimated from different ESMs
- 2. Assess AI-FDT projections of greenhouse gas and stratospheric aerosol injection (SAI) forcing
- 3. Optimizing MCB or SAI forcing patterns to produce specific climate responses in target regions (e.g., key regions for tipping points)

AI FDT projection procedure

- For each time lag τ :
 - Select N sample radiation anomaly fields $\vec{x_i}$ from a preprocessed dataset
 - Run AI (A_{τ}) on the N input field to obtain N control projected output fields $\overrightarrow{y_i}$ for each lag τ
 - Perturb each of the input radiation anomaly fields by $\delta \vec{f}$, $\vec{x'_i} = \vec{x_i} + \delta \vec{f}$
 - Run AI on the N perturbed input fields to obtain N perturbed output fields $\overrightarrow{y_i}$ for each lag τ
 - Compute average across the N control and perturbed output fields
- Compute Simpson's integration over the time lags au at each grid point



piAnthro effect



Global Mean Al lagged responses

- AI model suggests a response time scale to the MCB perturbation of 3-years
- Temperature response has not converged to 0 after 60 months, possibly suggesting longer time scales are required.



AI Emulation performance

- To test the AI performance, we:
 - Run AI model with CESM2 internal variability as input
 - Compare AI model output (tas, pr, ps) to the original CESM2 variables at the corresponding lag
- The AI model performs well in low latitude oceanic regions, but poorly at high latitudes and over land



Al Emulation performance

- AI model emulation has skill above noise out to ~36 months
- Substantially higher skill in spatial correlation for precipitation than surface temperature/pressure



Comparison between AI models

